From closed-boundary to single-sided homogeneous Green’s function representations
Kees Wapenaar*, Joost van der Neut, Jan Thorbecke and Evert Slob, Delft University of Technology; Satyan Singh, Colorado School of Mines

SUMMARY

The homogeneous Green’s function (i.e., the Green’s function and its time-reversal counterpart) plays an important role in optical, acoustic and seismic holography, in inverse scattering methods, in the field of time-reversal acoustics, in reverse-time migration and in seismic interferometry. Starting with the classical closed-boundary representation of the homogeneous Green’s function, we modify the configuration to two parallel boundaries. We discuss step-by-step a process that eliminates the integral along the lower boundary. This leads to a single-sided representation of the homogeneous Green’s function. Apart from imaging, we foresee interesting applications in inverse scattering, time-reversal acoustics, seismic interferometry, passive source imaging, etc.

INTRODUCTION

The homogeneous Green’s function plays an important role in optical, acoustic and seismic holography (Porter, 1970; Maynard et al., 1985; Wu and Toksöz, 1987; Lindsey and Braun, 2004), in linear inverse source problems and inverse scattering methods (Porter and Devaney, 1982; Oristaglio, 1989), in the field of time-reversal acoustics (Fink, 1997, 2008), in reverse-time migration (McMechan, 1983; Esmeroy and Oristaglio, 1988) and in seismic interferometry (Wapenaar, 2003; Derode et al., 2003; Weaver and Lobkis, 2004). The homogeneous Green’s function is formed by a combination of the causal Green’s function and its time-reversed version. An exact integral representation exists, but it is expressed in terms of a closed boundary integral. Here we explain in detail with numerical examples how the closed boundary integral can be transformed into an open integral, which thus leads to a single-sided integral representation of the homogeneous Green’s function (Wapenaar et al., 2016). This single-sided representation has interesting applications in the fields mentioned above.

THE HOMOGENEOUS GREEN’S FUNCTION

Consider a Green’s function \( G(x, x_B, t) \), defined as the response to an impulsive source of volume injection rate at \( x_B \). It is the causal solution of the acoustic wave equation with a source term \(-\delta(x-x_B)\delta(t)\) on the right-hand side. For a lossless medium the wave equation is symmetric in time, except for the source term, which is antisymmetric in time. Hence, the time-reversed Green’s function \( G(x, x_B, -t) \) obeys the same wave equation, but with an opposite source term \(+\delta(x-x_B)\delta(t)\). The sum of the Green’s function and its time-reversal, i.e., \( G_h(x, x_B, t) = G(x, x_B, t) + G(x, x_B, -t) \) also obeys the same wave equation, but with the source terms cancelling each other. Since in this case the right-hand side of the wave equation is zero, we speak of a homogeneous equation, and we call its solution \( G_h(x, x_B, t) \) the homogeneous Green’s function (not to be confused with the Green’s function for a homogeneous medium). In the frequency domain, a representation for the homogeneous Green’s function reads (Porter, 1970; Oristaglio, 1989; Wapenaar et al., 2005)

\[
G_h(x_A, x_B, \omega) = \frac{1}{j \omega \rho} \int_{\partial D} \{ G^1(x, x_A, \omega) \partial_t \tilde{G}(x, x_B, \omega) - \partial_t \tilde{G}^*(x, x_A, \omega) G(x, x_B, \omega) \} \eta_i d^2x. \tag{1}
\]

where \( \omega \) denotes angular frequency, \( \rho \) mass density, \( j \) the imaginary unit and \( * \) complex conjugation. \( \partial \mathbb{D} \) is a closed boundary with outward pointing normal vector \( (n_1, n_2, n_3) \), enclosing a domain \( D \), and \( x_A \) and \( x_B \) are the coordinate vectors of two points inside \( \partial \mathbb{D} \), see Figure 1(a). Equation (1) is exact and thus accounts for all orders of multiple scattering inside and outside domain \( \mathbb{D} \).

For the configuration of Figure 1(b), we modify equation (1) as follows

\[
G(x_A, x_B, \omega) + \tilde{G}^*(x_B, x_A, \omega) \tag{2}
\]

\[
= \int_{\partial D_A} \frac{1}{j \omega \rho} \{ \tilde{G}_{3A} \partial_t G_B - \partial_t \tilde{G}_A^* G_B \} d^2x.
\]

\[
- \int_{\partial D_C} \frac{1}{j \omega \rho} \{ \tilde{G}_{3A} \partial_t G_B - \partial_t \tilde{G}_A^* G_B \} d^2x.
\]
Homogeneous Green’s function representation

where $\partial \mathcal{D}_R$ and $\partial \mathcal{D}_C$ are two infinite horizontal boundaries. The contribution of the integral along the cylindrical boundary $\partial \mathcal{D}_{cyl}$ vanishes. $\hat{G}$ and $G$ are short-hand notations for $\hat{G}(x, x_A, \omega)$ and $G(x, x_B, \omega)$, respectively. We replaced $G_A$ by a Green’s function $\bar{G}$ in a reference medium, which is identical to the actual medium below $\partial \mathcal{D}_R$, but homogeneous at and above $\partial \mathcal{D}_R$. To arrive at a single-sided integral representation, we have to eliminate the integral along the lower boundary $\partial \mathcal{D}_C$. This is the subject of the next section.

AN AUXILIARY FUNCTION

We introduce an auxiliary function $\Gamma(x, \omega)$ which we subtract from the Green’s function, according to

$$\hat{G}(x, x_A, \omega) \rightarrow \hat{G}(x, x_A, \omega) - \Gamma(x, \omega). \quad (3)$$

As long as $\Gamma(x, \omega)$ obeys the same wave equation in $\bar{D}$ as $\hat{G}(x, x_A, \omega)$, but without a source term, we can make this replacement in equation (2), hence,

$$G(x_A, x_B, \omega) + \{G(x_B, x_A, \omega) - \Gamma(x_B, \omega)\}^*$$

$$= \int_{\partial \mathcal{D}_A} \frac{1}{16\pi} \left\{ (G_A - \Gamma)^* \partial_3 G_B - \partial_3 (G_A - \Gamma)^* G_B \right\} d^2 x$$

$$- \int_{\partial \mathcal{D}_C} \frac{1}{16\pi} \left\{ (\bar{G}_A - \Gamma)^* \partial_3 G_B - \partial_3 (\bar{G}_A - \Gamma)^* G_B \right\} d^2 x. \quad (4)$$

We search for a function $\Gamma$, such that both $G_A - \Gamma$ and $\partial_3 (G_A - \Gamma)$ vanish on $\partial \mathcal{D}_C$ (these are the Dirichlet and Neumann boundary conditions, which together are known as the Cauchy boundary condition, hence the subscript C in $\partial \mathcal{D}_C$). Introducing auxiliary functions is a common approach to manipulate the boundary conditions (Morse and Feshbach, 1953; Berkhout, 1982). For the integral in equation (4) this was previously proposed by Weglein et al. (2011), but solved only for some special cases. Recently we proposed a more general way to find a $\Gamma$ that obeys both boundary conditions for arbitrary inhomogeneous media (Wapenaar et al., 2016). Here we explain this method in more detail and illustrate it step by step with a numerical example. Although the numerical example is 1D, the proposed approach holds for 3D inhomogeneous media.

We consider a horizontally layered medium, with interfaces at $z = 300, 600$ and $900$ m. The propagation velocities in the layers are $1500, 1950, 2000$ and $2300$ m/s and the mass densities are $1000, 4500, 1400$ and $1600$ kg/m$^3$, respectively. A Green’s source is defined at $z_A = 800$ m. The 1D time-domain Green’s function $G(z, z_A, t)$ is shown in a VSP-like display in Figure 2. The red dot denotes the source, the red lines the direct arrivals. The traces in the top and bottom panels are the responses at $z_R = 0$ and $z_C = 1175$m, respectively, denoted by the blue dots. The auxiliary function $\Gamma(z, t)$ should be defined such that at $z_C$ it cancels the Green’s function $G(z, z_A, t)$ (i.e., the trace in the bottom panel of Figure 2). The focusing functions, introduced earlier for Marchenko imaging (Broggini and Snieder, 2012; Wapenaar et al., 2013; Slob et al., 2014; van der Neut et al., 2015), can generate such a function. The top panel in Figure 3 shows the focusing function $f^+_1(z_R, z_A, t)$, which is emitted from $z_R = 0$ into the medium. The VSP-like panel shows the evolution of this focusing function through the medium (left of the red lines). The field focuses at $z_A = 800$ m (the yellow dot). The focused field at $z_A$ acts as a virtual source for downgoing waves, of which the response is denoted as $G^{p, +}(z, z_A, t)$ (where the second superscript, +, denotes the downgoing source at $z_A$, and the first superscript, p, the total pressure field at $z$). This response is shown right of the red lines in Figure 3. The response at $z_C$, denoted by the blue dot, is shown in the lower panel of Figure 3. It contains part of the events of $G(z_C, z_A, t)$ in Figure 2. The events still missing in Figure 3 are those caused by the upward radiating part of the source at the red dot in Figure 2. We now discuss how this remaining part of $G(z_C, z_A, t)$ can be recovered by another focusing function. Consider again the focusing function $f^+_1(z, z_A, t)$ in Figure 3. Before reaching the focus, a part of this focusing function is reflected upward and is called $f^-_1(z, z_A, t)$. At $z = z_R$ we reverse this field in time and change its polarity, yielding $-f^-_1(z_R, z_A, t)$. Figure 4 shows the emission of this new focusing function into the medium. Left of the red lines, its response is $-f^-_1(z, z_A, t) - f^+_1(z, z_A, t)$. The response right of the red lines apparently originates from a source for upgoing waves at the yellow dot at $z_A$, hence, this response is denoted as $G^{p, -}(z, z_A, t)$ (where the second superscript, -, denotes the upgoing source at $z_A$). The trace in the lower panel in Figure 4 shows the events of $G(z_C, z_A, t)$ (Figure 2) that were missing.
Homogeneous Green’s function representation

in Figure 3. The superposition of Figures 3 and 4 constitutes the desired auxiliary function \( \Gamma(z,t) \), because at \( z_C \) this gives

\[
\Gamma(z,t) = \mathcal{G}_{z_C,z_A,t}^{+} + \mathcal{G}_{z_C,z_A,t}^{-} = \mathcal{G}(z_C,z_A,t),
\]

see Figure 5. Left of the red lines (and above the yellow dot) the field consists of \( \mathcal{H}(z_a - z)\left( f_1(z,z_A,t) - f_1(z,z_A,-t) \right) \), where \( \mathcal{H}(z) \) is the Heaviside function and \( f_1(z,z_A,t) = f_1^+(z,z_A,t) + f_1^-(z,z_A,t) \), and right of the red lines it is \( \mathcal{G}(z,z_A,t) \). Together, this gives

\[
\Gamma(z,t) = \mathcal{G}(z,z_A,t) + \mathcal{H}(z_a - z)\left( f_1(z,z_A,t) - f_1(z,z_A,-t) \right).
\]

Hence, by subtracting \( \Gamma(z,t) \) from \( \mathcal{G}(z,z_A,t) \), i.e., \( \mathcal{G}(z,z_A,t) - \Gamma(z,t) \) (Figure 6), we are left with the focusing function and its time-reversal above \( z_A \) (the yellow dot). The field in the half-space below the yellow dot is zero, hence also its vertical derivative is zero, so the Dirichlet and Neumann boundary conditions are both obeyed at \( z = z_C \).

SINGLE-SIDED REPRESENTATIONS

Following a more formal 3D derivation in the space-frequency domain, we obtain analogous to (5)

\[
\Gamma(x,\omega) = \mathcal{G}(x,x_A,\omega) + \mathcal{H}(z_a - z)2\mathcal{I}\{ f_1(x,x_A,\omega) \},
\]

where \( \mathcal{I} \) denotes the imaginary part (Wapenaar et al., 2016).

Substitution into equation (4), taking the real part of both sides, and using \( 2\mathcal{R}\{ \mathcal{G}(x,x_B,\omega) \} = \mathcal{G}_h(x,x_B,\omega) \), gives

\[
\mathcal{G}_h(x_A,x_B,\omega) = \int_{\mathbb{D}_B} \frac{2}{\mathcal{D}(\mathbf{x})} \left( \mathcal{I}\{ f_1(x,x_A,\omega) \} \partial_3 \mathcal{G}_h(x,x_B,\omega) - \mathcal{G}_h(x,x_B,\omega) \right) \mathcal{D}^2 \mathbf{x},
\]

see Figure 7. Note that the Green’s function \( \mathcal{G}_h(x,x_B,\omega) \) under the integral can be obtained from a similar representation. With some simple replacements (and using source-receiver reciprocity) we obtain

\[
\mathcal{G}_h(x,x_B,\omega) = \int_{\mathbb{D}_B} \frac{2}{\mathcal{D}(\mathbf{x})} \left( \mathcal{I}\{ f_1(x',x_B,\omega) \} \partial_3 \mathcal{G}_h(x',x_B,\omega) - \mathcal{G}_h(x',x_B,\omega) \right) \mathcal{D}^2 \mathbf{x}',
\]

with \( x \) on \( \partial \mathbb{D}_B \) and \( x' \) on \( \partial \mathbb{D}_C \), just above \( \partial \mathbb{D}_B \). Note that \( \mathcal{G}_h(x',\omega) \) stands for the reflection response at the surface. Hence, equations (7) and (8) can be used to retrieve the homogeneous Green’s function \( \mathcal{G}_h(x_A,x_B,\omega) \) from the reflection response \( \mathcal{G}_h(x',\omega) \). This two-step process is summarised as

\[
\mathcal{G}_h(x',\omega) \xrightarrow{\text{focusing}} \mathcal{G}_h(x,x_B,\omega) \xrightarrow{\text{reflection}} \mathcal{G}_h(x_A,x_B,\omega).
\]
Homogeneous Green’s function representation

![Homogeneous Green’s function representation](image)

**CONCLUSIONS**

Starting with the classical homogeneous Green’s function representation for the configuration of Figure 1(a) (equation 1), we modified the configuration to two parallel boundaries ∂DR and ∂DC (Figure 1(b)), and discussed a way to eliminate the integral along the lower boundary ∂DC. To this end we introduced an auxiliary function, which consists of focusing functions, emitted from the upper boundary, which reproduce the Green’s function at the lower boundary. Hence, by subtracting this auxiliary function from the Green’s function, the integral along the lower boundary vanishes, leaving a single-sided representation of the homogeneous Green’s function (Figure 7).

Note that the focusing functions appearing in the single-sided representation are those we derived earlier for Marchenko imaging. These focusing functions can be retrieved from the reflection response at the surface and an estimate of the direct arrival between the focal point and the surface. Hence, the two-step redatuming process, summarised by equation (9), handles multiple reflections in a data-driven way. Apart from imaging, we foresee interesting applications of the single-sided homogeneous Green’s function representation in inverse scattering, time-reversal acoustics, seismic interferometry, passive source imaging, etc.

![Visualisation of the single-sided homogeneous Green’s function representation](image)
REFERENCES